

MOX-Report No. 68/2022

An IMEX-DG solver for atmospheric dynamics simulations with adaptive mesh refinement

Orlando, G.; Benacchio, T.; Bonaventura, L.

MOX, Dipartimento di Matematica Politecnico di Milano, Via Bonardi 9 - 20133 Milano (Italy)

mox-dmat@polimi.it

http://mox.polimi.it

An IMEX-DG solver for atmospheric dynamics simulations with adaptive mesh refinement

Giuseppe Orlando⁽¹⁾ Tommaso Benacchio⁽²⁾, Luca Bonaventura⁽¹⁾

 (1) MOX, Dipartimento di Matematica, Politecnico di Milano Piazza Leonardo da Vinci 32, 20133 Milano, Italy giuseppe.orlando@polimi.it, luca.bonaventura@polimi.it

> ⁽²⁾ Future Rotorcraft Technologies, Leonardo S.p.a. Cascina Costa di Samarate (VA), Italy tommaso.benacchio.ext@leonardo.com

Keywords: Numerical weather prediction, Adaptive mesh refinement, Discontinuous Galerkin methods, Flows over orography.

AMS Subject Classification: 65M12, 65M50, 65M60, 65Y05, 76N06

Abstract

We present an accurate and efficient solver for atmospheric dynamics simulations that allows for non-conforming mesh refinement. The model equations are the conservative Euler equations for compressible flows. The numerical method is based on an h-adaptive Discontinuous Galerkin spatial discretization and on a second order Additive Runge Kutta IMEX method for time discretization, especially designed for low Mach regimes. The solver is implemented in the framework of the *deal.II* library, whose mesh refinement capabilities are employed to enhance efficiency. A number of numerical experiments based on classical benchmarks for atmosphere dynamics demonstrate the properties and advantages of the proposed method.

1 Introduction

Atmospheric dynamics can be described by mathematical models tailored for specific scales, but use of the fully compressible equations has become the standard in the last couple of decades, see e.g. the literature review in [31]. The efficient numerical solution of these equations in presence of gravitational forcing is crucial for weather and climate predictions and poses several computational challenges.

The Discontinuous Galerkin (DG) discretization has proven itself a valuable tool in environmental fluid dynamics models, see e.g. [14, 15] or the review in [5]. Discontinuous finite element h-adaptive approaches were proposed in [19, 22, 35], while p-adaptive DG methods for Numerical Weather Prediction (NWP) were introduced in [34]. However, severe time step stability restrictions may be required if the DG method is coupled with explicit time discretization schemes. The use of implicit and semi-implicit methods, which has a long tradition in the literature on atmosphere dynamics and numerical weather prediction, see among many others [3, 4, 11, 12, 21, 30], allows one to relax these restrictions and employ much longer time steps.

In this work, we will present an application of the h-adaptive DG method recently proposed in [23] to a number of idealized benchmarks for atmospheric flow modelling. The proposed numerical approach combines accurate and flexible DG space discretization with an implicit-explicit (IMEX) time discretization, whose properties and theoretical justifications are discussed in detail in [23] and briefly recalled in the following. The adaptive discretization is implemented in the framework of the numerical library deal. II [1, 2], an open-source environment, which provides non-conforming h-refinement capabilities that will be exploited in the numerical simulation of atmospheric flows. This choice allows us to embed the development of codes specific for numerical weather prediction applications into a framework that is being continuously adapted to novel architectures and extended to include novel versions of the DG method. As a result, the portability and the continuous development of the resulting code are greatly enhanced. The accuracy of the results obtained and the performance achieved by the model show that a dynamical core for high resolution, fine scale atmospheric modelling can be based on a widely used open source finite element library. It will also be shown that the proposed adaptive implementation can achieve good parallel efficiency, even without employing specific optimizations.

The paper is structured as follows. The model equations are reviewed in Section 2. The time and space discretizations are briefly outlined and discussed in Section 3. The validation of the proposed method and its applications to a number of significant benchmarks with and without orography are reported in Sections 4, 5, respectively. Some conclusions and considerations about open issues and future work are presented in Section 6.

2 The model equation

Let $\Omega \subset \mathbb{R}^d, 2 \leq d \leq 3$ be a connected open bounded set with a sufficiently smooth boundary $\partial \Omega$ and denote by **x** the spatial coordinates and by t the temporal coordinate. We consider the classical unsteady compressible Euler equations, written in flux form as:

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{u}) = 0$$

$$\frac{\partial (\rho \mathbf{u})}{\partial t} + \nabla \cdot (\rho \mathbf{u} \otimes \mathbf{u}) + \nabla p = -\rho g \mathbf{k} \qquad (1)$$

$$\frac{\partial (\rho E)}{\partial t} + \nabla \cdot [(\rho E + p) \mathbf{u}] = -\rho g \mathbf{k} \cdot \mathbf{u}$$

for $\mathbf{x} \in \Omega$, $t \in [0, T_f]$, supplied with suitable initial and boundary conditions. Here T_f is the final time, ρ is the density, \mathbf{u} is the fluid velocity, p is the pressure, $g = 9.81 \,\mathrm{m \, s^{-2}}$ is the acceleration of gravity and \mathbf{k} is the upward pointing unit vector in the standard Cartesian reference frame. The total energy ρE can be rewritten as $\rho E = \rho e + \rho k$, where e is the internal energy and $k = \frac{1}{2} ||\mathbf{u}||^2$ is the kinetic energy. We also introduce the specific enthalpy $h = e + \frac{p}{\rho}$ and we notice that one can rewrite the energy flux as

$$(\rho E + p)\mathbf{u} = \left(e + k + \frac{p}{\rho}\right)\rho\mathbf{u} = (h+k)\rho\mathbf{u}.$$

The equations can then be rewritten as

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{u}) = 0$$

$$\frac{\partial (\rho \mathbf{u})}{\partial t} + \nabla \cdot (\rho \mathbf{u} \otimes \mathbf{u}) + \nabla p = -\rho g \mathbf{k}$$

$$\frac{\partial (\rho E)}{\partial t} + \nabla \cdot [(h+k)\rho \mathbf{u}] = -\rho g \mathbf{k} \cdot \mathbf{u}.$$
(2)

The above equations are complemented by the equation of state (EOS) for ideal gases, given by $p = \rho RT$, with R being the specific gas constant and T the absolute temperature. As explained in [23], the proposed method also allows to handle in the same framework also more general equations of state for real gases, thus opening the way for example to the inclusion of effects due to water vapour and moist species. Furthermore, while we describe here for simplicity only the numerical approximation of the Euler equations, as customary when describing dynamical cores for atmospheric applications, it was shown again in [23] how to extend the proposed approach to the viscous case by Strang splitting [33]. This extension will be required by some of the numerical simulations described in Section 4.

We also introduce the potential temperature θ , which is often preferred in atmospheric applications as the unknown in the energy equation. For an ideal gas, potential temperature is defined as

$$\theta = T\left(\frac{p_0}{p}\right)^{\frac{\gamma-1}{\gamma}},$$

where p_0 is a reference pressure and γ denotes the specific heats ratio. In the present work, we consider $p_0 = 10^5$ Pa.

3 The numerical framework

In the low Mach number limit, pressure gradient terms yield stiff components for the resulting semidiscretized ODE system. Therefore, following [7, 8], the method proposed in [23] couples implicitly the energy equation to the momentum one, while the continuity equation is treated in a fully explicit fashion. For the time discretization, an IMplicit EXplicit (IMEX) Additive Runge Kutta method (ARK) [16] method will be used. These methods are widely employed for ODE systems that include both stiff and non-stiff components, to which the implicit and explicit schemes are applied, respectively. Here, we consider a variant of the IMEX method proposed in [12], whose coefficients are presented in the Butcher tableaux reported in Tables 1 and 2 for the explicit and implicit method, respectively.

$$\begin{array}{cccccc} 0 & 0 & & \\ \chi & \chi & 0 & \\ 1 & 1 - \alpha & \alpha & 0 \\ \hline & \frac{1}{2} - \frac{\chi}{4} & \frac{1}{2} - \frac{\chi}{4} & \frac{\chi}{2} \end{array}$$

Table 1: Butcher tableau of the explicit ARK2 method

We consider $\chi = 2 - \sqrt{2}$, so that the implicit part of the IMEX scheme coincides with the TR-BDF2 method considered in [34]. A solver based on this method for the incompressible



Table 2: Butcher tableau of the implicit ARK2 method

Navier Stokes equations with an artificial compressibility formulation has been proposed in [24], highlighting the robustness of the proposed approach in the low Mach number limit. Notice also that, as discussed in [23], we take $\alpha = 0.5$, rather than the value originally chosen in [12], in order to improve the monotonicity of the method.

We now briefly describe the application of this IMEX method outlined above to equations (2). We present for simplicity only the time semi-discretization, while the reader is referred to [23] for the description of the fully discrete scheme and the detailed formulation of the algebraic equations involving the discrete degrees of freedom. As it can be seen from the Butcher tableaux, the first stage is only formal, yielding $\rho^{(n,1)} = \rho^n$ for density and all the other prognostic quantities. The second stage reads as follows

$$\rho^{(n,2)} = \rho^{n} - a_{21} \Delta t \, \nabla \cdot (\rho^{n} \mathbf{u}^{n})
\rho^{(n,2)} \mathbf{u}^{(n,2)} + \tilde{a}_{22} \Delta t \nabla p^{(n,2)} = \mathbf{m}^{(n,2)}
\rho^{(n,2)} E^{(n,2)} + \tilde{a}_{22} \Delta t \, \nabla \cdot \left(h^{(n,2)} \rho^{(n,2)} \mathbf{u}^{(n,2)} \right) = \mathbf{e}^{(n,2)},$$
(3)

where we have set

$$\mathbf{m}^{(n,2)} = \rho^{n} \mathbf{u}^{n} - a_{21} \Delta t \, \nabla \cdot (\rho^{n} \mathbf{u}^{n} \otimes \mathbf{u}^{n}) - \tilde{a}_{21} \Delta t \nabla p^{n} - \tilde{a}_{21} \Delta t \rho^{n} g \mathbf{k} - \tilde{a}_{22} \Delta t \rho^{(n,2)} g \mathbf{k}$$
(4)

$$\mathbf{e}^{(n,2)} = \rho^{n} E^{n} - \tilde{a}_{21} \Delta t \, \nabla \cdot (h^{n} \rho^{n} \mathbf{u}^{n}) - a_{21} \Delta t \, \nabla \cdot (k^{n} \rho^{n} \mathbf{u}^{n}) - \tilde{a}_{21} \rho^{n} g \mathbf{k} \cdot \mathbf{u}^{n} - \tilde{a}_{22} \rho^{(n,2)} g \mathbf{k} \cdot \mathbf{u}^{(n,2)}.$$

For the third stage, one can write formally

$$\rho^{(n,3)} = \rho^{n} - a_{31}\Delta t \,\nabla \cdot \left(\rho^{n}\mathbf{u}^{n}\right) - a_{32}\Delta t \,\nabla \cdot \left(\rho^{(n,2)}\mathbf{u}^{(n,2)}\right)
\rho^{(n,3)}\mathbf{u}^{(n,3)} + \tilde{a}_{33}\Delta t \nabla p^{(n,3)} = \mathbf{m}^{(n,3)}
\rho^{(n,3)}E^{(n,3)} + \tilde{a}_{33}\Delta t \,\nabla \cdot \left(h^{(n,3)}\rho^{(n,3)}\mathbf{u}^{(n,3)}\right) = \mathbf{e}^{(n,3)},$$
(5)

where the right hand sides are defined as

$$\mathbf{m}^{(n,3)} = \rho^{n} \mathbf{u}^{n}$$

$$- a_{31} \Delta t \,\nabla \cdot \left(\rho^{n} \mathbf{u}^{n} \otimes \mathbf{u}^{n}\right) - \tilde{a}_{31} \Delta t \nabla p^{n}$$

$$- a_{32} \Delta t \,\nabla \cdot \left(\rho^{(n,2)} \mathbf{u}^{(n,2)} \otimes \mathbf{u}^{(n,2)}\right) - \tilde{a}_{32} \Delta t \nabla p^{(n,2)}$$

$$- \tilde{a}_{31} \Delta t g \rho^{n} \mathbf{k} - \tilde{a}_{32} \Delta t g \rho^{(n,2)} \mathbf{k} - \tilde{a}_{33} \Delta t g \rho^{(n,3)} \mathbf{k}$$
(6)

$$\mathbf{e}^{(n,3)} = \rho^{n} E^{n} - \tilde{a}_{31} \Delta t \,\nabla \cdot (h^{n} \rho^{n} \mathbf{u}^{n}) - a_{31} \Delta t \,\nabla \cdot (k^{n} \rho^{n} \mathbf{u}^{n}) - \tilde{a}_{31} \Delta t \rho^{n} g \mathbf{k} \cdot \mathbf{u}^{n} - \tilde{a}_{32} \Delta t \,\nabla \cdot \left(h^{(n,2)} \rho^{(n,2)} \mathbf{u}^{(n,2)}\right) - a_{32} \Delta t \,\nabla \cdot \left(k^{(n,2)} \rho^{(n,2)} \mathbf{u}^{(n,2)}\right) - \tilde{a}_{32} \Delta t \rho^{(n,2)} g \mathbf{k} \cdot \mathbf{u}^{(n,2)} - \tilde{a}_{33} \Delta t \rho^{(n,3)} g \mathbf{k} \cdot \mathbf{u}^{(n,3)}.$$

In each stage, after spatial discretization has been performed, we formally substitute the momentum into the energy equation and we obtain an equation for the pressure [23].

For what concerns the spatial discretization, we consider a decomposition of the domain Ω into a family of hexahedra \mathcal{T}_h (quadrilaterals in the two-dimensional case) and denote each element by K. The skeleton \mathcal{E} denotes the set of all element faces and $\mathcal{E} = \mathcal{E}^I \cup \mathcal{E}^B$, where \mathcal{E}^I and \mathcal{E}^B are the subset of interior and boundary faces, respectively. Suitable jump and average operators can then be defined as customary for finite element discretizations. A face $\Gamma \in \mathcal{E}^I$ shares two elements that we denote by K^+ with outward unit normal \mathbf{n}^+ and K^- with outward unit normal \mathbf{n}^- , whereas for a face $\Gamma \in \mathcal{E}^B$ we simply denote by \mathbf{n} the outward unit normal. For a scalar function φ the jump is defined as

$$[[\varphi]] = \varphi^{+} \mathbf{n}^{+} + \varphi^{-} \mathbf{n}^{-} \quad \text{if } \Gamma \in \mathcal{E}^{I} \qquad [[\varphi]] = \varphi \mathbf{n} \quad \text{if } \Gamma \in \mathcal{E}^{B}.$$

The average is defined as

$$\{\{\varphi\}\} = \frac{1}{2} \left(\varphi^+ + \varphi^-\right) \quad \text{if } \Gamma \in \mathcal{E}^I \qquad \{\{\varphi\}\} = \varphi \quad \text{if } \Gamma \in \mathcal{E}^B.$$

Similar definitions apply for a vector function φ :

$$\begin{split} & [[\boldsymbol{\varphi}]] = \boldsymbol{\varphi}^+ \cdot \mathbf{n}^+ + \boldsymbol{\varphi}^- \cdot \mathbf{n}^- \quad \text{if } \Gamma \in \mathcal{E}^I \qquad [[\boldsymbol{\varphi}]] = \boldsymbol{\varphi} \cdot \mathbf{n} \quad \text{if } \Gamma \in \mathcal{E}^B \\ & \{\{\boldsymbol{\varphi}\}\} = \frac{1}{2} \left(\boldsymbol{\varphi}^+ + \boldsymbol{\varphi}^-\right) \quad \text{if } \Gamma \in \mathcal{E}^I \qquad \{\{\boldsymbol{\varphi}\}\} = \boldsymbol{\varphi} \quad \text{if } \Gamma \in \mathcal{E}^B. \end{split}$$

For vector functions, it is also useful to define a tensor jump as:

$$\langle\langle \boldsymbol{\varphi} \rangle \rangle = \boldsymbol{\varphi}^+ \otimes \mathbf{n}^+ + \boldsymbol{\varphi}^- \otimes \mathbf{n}^- \quad \text{if } \Gamma \in \mathcal{E}^I \qquad \langle\langle \boldsymbol{\varphi} \rangle \rangle = \boldsymbol{\varphi} \otimes \mathbf{n} \quad \text{if } \Gamma \in \mathcal{E}^B.$$

We introduce the following finite element spaces

$$Q_r = \left\{ v \in L^2(\Omega) : v |_K \in \mathbb{Q}_r \quad \forall K \in \mathcal{T}_h \right\} \quad \mathbf{V}_r = [Q_r]^d,$$

where \mathbb{Q}_r is the space of polynomials of degree r in each coordinate direction. We then denote by $\varphi_i(\mathbf{x})$ the basis functions for the space \mathbf{V}_r and by $\psi_i(\mathbf{x})$ the basis functions for the space Q_r , the finite element spaces chosen for the discretization of the velocity and of the pressure (as well as the density), respectively, so that

$$\mathbf{u} \approx \sum_{j=1}^{\dim(\mathbf{V}_r)} u_j(t) \boldsymbol{\varphi}_j(\mathbf{x}) \qquad p \approx \sum_{j=1}^{\dim(Q_r)} p_j(t) \psi_j(\mathbf{x}).$$

The shape functions are chosen to be the Lagrange interpolation polynomials for the support points of (r+1)-order Gauss-Lobatto quadrature rule along each coordinate direction. Given these definitions, the weak formulation for the momentum equation at each stage s = 2, 3 of the IMEX scheme can be written in compact form as

$$\mathbf{A}^{(n,s)}\mathbf{U}^{(n,s)} + \mathbf{B}^{(n,s)}\mathbf{P}^{(n,s)} = \mathbf{F}^{(n,s)},\tag{7}$$

where we have set

$$A_{ij}^{(n,s)} = \sum_{K} \int_{K} \rho^{(n,s)} \varphi_j \cdot \varphi_i d\Omega$$
(8)

$$B_{ij}^{(n,s)} = \sum_{K} \int_{K} -\tilde{a}_{ss} \Delta t \nabla \cdot \boldsymbol{\varphi}_{i} \psi_{j} d\Omega + \sum_{\Gamma \in \mathcal{E}} \int_{\Gamma} \tilde{a}_{ss} \Delta t \left\{ \{\psi_{j}\} \} \left[[\boldsymbol{\varphi}_{i}] \right] d\Sigma$$

$$(9)$$

with $\mathbf{U}^{(n,s)}$ denoting the vector of the degrees of freedom associated to the velocity field and $\mathbf{P}^{(n,s)}$ denoting the vector of the degrees of freedom associated to the pressure. Moreover,

 $\rho^{(n,s)}$ denotes the approximation of the density at stage s and \tilde{a}_{ss} is the corresponding coefficient of the Butcher tableaux Table 2 for the implicit part. Notice that the definitions introduced above entail that centered fluxes are used in the definition of the discrete pressure gradient, while an upwind flux has been employed in the definition of $\mathbf{F}^{(n,s)}$ for the discretization of the terms computed at time level n. Analogously, the weak formulation for the energy equation can be expressed as

$$\mathbf{C}^{(n,s)}\mathbf{U}^{(n,s)} = \mathbf{G}^{(n,s)},\tag{10}$$

where we have set

$$C_{ij}^{(n,s)} = \sum_{K} \int_{K} -\tilde{a}_{ss} \Delta t h^{(n,s)} \rho^{(n,s)} \varphi_{j} \cdot \nabla \psi_{i} d\Omega + \sum_{\Gamma \in \mathcal{E}} \int_{\Gamma} \tilde{a}_{ss} \Delta t \left\{ \left\{ h^{(n,s)} \rho^{(n,s)} \varphi_{j} \right\} \right\} \cdot [[\psi_{i}]] d\Sigma,$$
(11)

where $h^{(n,s)}$ represents the approximation of the enthalpy at stage *s*. Formally, we can then derive $\mathbf{U}^{(n,s)} = (\mathbf{A}^{(n,s)})^{-1} (\mathbf{F}^{(n,s)} - \mathbf{B}^{(n,s)} \mathbf{P}^{(n,s)})$ and obtain the following relation

$$\mathbf{C}^{(n,s)}(\mathbf{A}^{(n,s)})^{-1}\left(\mathbf{F}^{(n,s)} - \mathbf{B}^{(n,s)}\mathbf{P}^{(n,s)}\right) = \mathbf{G}^{(n,s)}.$$
(12)

Again, in the definition of $\mathbf{G}^{(n,s)}$ an upwind flux has been employed for the discretization of the terms computed at time level n. Moreover, taking into account that

$$\rho^{(n,s)}E^{(n,s)} = \rho^{(n,s)}e^{(n,s)}(p^{(n,s)}) + \rho^{(n,s)}k^{(n,s)},$$

we decompose $\mathbf{G}^{(n,s)}$ and we finally obtain

$$\mathbf{C}^{(n,s)}(\mathbf{A}^{(n,2)})^{-1}\left(\mathbf{F}^{(n,s)} - \mathbf{B}^{(n,s)}\mathbf{P}^{(n,s)}\right) = -\mathbf{D}^{(n,s)}\mathbf{P}^{(n,2)} + \tilde{\mathbf{G}}^{(n,s)},$$
(13)

where we have set

$$D_{ij}^{(n,s)} = \sum_{K} \int_{K} \rho^{(n,s)} e^{(n,s)} (\psi_j) \psi_i d\Omega.$$
 (14)

The above system can be solved in terms of $\mathbf{P}^{(n,s)}$ according to the fixed point procedure described in [8]. More specifically, setting

$$\mathbf{P}^{(n,s,0)} = \mathbf{P}^{(n,s-1)}, k^{(n,s,0)} = k^{(n,s-1)},$$

for $l = 1, \ldots, M$ one solves the equation

$$\begin{pmatrix} \mathbf{D}^{(n,s,l)} - \mathbf{C}^{(n,s,l)} (\mathbf{A}^{(n,s)})^{-1} \mathbf{B}^{(n,s)} \end{pmatrix} \mathbf{P}^{(n,s,l+1)} = \\ \tilde{\mathbf{G}}^{(n,s,l)} - \mathbf{C}^{(n,s,l)} (\mathbf{A}^{(n,s)})^{-1} \mathbf{F}^{(n,s,l)}$$

and updates the velocity solving

$$\mathbf{A}^{(n,s)}\mathbf{U}^{(n,s,l+1)} = \mathbf{F}^{(n,s,l)} - \mathbf{B}^{(n,s)}\mathbf{P}^{(n,s,l+1)}.$$

Once the iterations have been completed, one sets $\mathbf{u}^{(n,s)} = \mathbf{u}^{(n,s,M+1)}$ and $E^{(n,s)}$ accordingly. It is important to point out that the scheme outlined above only requires the solution of linear systems of a size equal to that of the number of discrete degrees of freedom associated to a scalar variable, as in [8], which is crucial for the overall efficiency.

4 Numerical results without orography

The numerical method outlined in the previous Sections has been validated in a number of benchmarks relevant for atmospheric applications. We define

$$\mathcal{H} = \min\{\operatorname{diam}(K) | K \in \mathcal{T}_h\}$$

and we define two Courant numbers, one based on the speed of sound denoted by C, the so-called acoustic Courant number, and one based on the local velocity of the flow, the so-called advective Courant number, denoted by C_u :

$$C = rc\Delta t/\mathcal{H}, \qquad C_u = ru\Delta t/\mathcal{H}$$
 (15)

where c is the magnitude of the speed of sound and u is the magnitude of the flow velocity. Notice that both the Courant numbers depend on the polynomial degree r. In general, we consider $\gamma = 1.4$ and $R = 287 \, \mathrm{J \, kg^{-1} \, K^{-1}}$ for all the simulations. In this Section, we will focus on the results obtained in benchmarks without orography, while we will discuss benchmarks with non trivial orographic profiles in the following Section.

4.1 Inertia-gravity waves

Inertia-gravity waves in a two-dimensional vertical section of the atmosphere constitute a classical benchmark for atmospheric flow models, see e.g. [4, 21, 28]. In particular, we set

$$\bar{\theta} = T_{ref} \exp\left(\frac{N^2 z}{g}\right) \tag{16}$$

where $N = 0.01 \,\mathrm{s}^{-1}$ denotes the buoyancy frequency and $T_{ref} = 300 \,\mathrm{K}$. The background density and pressure are defined as

$$\bar{p} = \exp\left\{1 - \frac{g^2}{N^2} \frac{\gamma - 1}{\gamma} \frac{\rho_{ref}}{p_{ref}} \left[1 - \exp\left(-\frac{N^2 z}{g}\right)\right]\right\}$$
(17)

$$\bar{\rho} = \rho_{ref} \left(\frac{p}{p_{ref}}\right)^{\frac{1}{\gamma}} \exp\left(-\frac{N^2 z}{g}\right)$$
(18)

with $p_{ref} = 10^5$ Pa and $\rho_{ref} = \frac{p_{ref}}{/} RT_{ref}$. The domain is $\Omega = (0, 300) \times (0, 10)$ km and we consider the following perturbation for the potential temperature

$$\theta' = 0.01 \frac{\sin\left(\frac{\pi y}{H}\right)}{1 + \left(\frac{x - x_c}{a}\right)^2} \tag{19}$$

with $x_c = 100 \text{ km}$, a = 5 km and H = 10 km. For what concerns the boundary conditions, we consider periodic boundary conditions for the horizontal direction and wall boundary conditions for the vertical direction. A background horizontal velocity $u = 20 \text{ m s}^{-1}$ is imposed. The grid is composed by 300×10 elements with r = 4, while the time step is taken equal to 3 s yielding $C \approx 4.17$ and $C_u \approx 0.24$. Figure 1 shows the the potential temperature perturbation at t = 3000 s, where one can easily notice that inertia-gravity waves propagate from the initial perturbation also reported in Figure 1. The results compare well with those available in the literature, see e.g. [21, 34]. Figure 2 shows the one-dimensional profile of the potential temperature perturbation along z = 5 km, which is symmetric about the position x = 160 km and in excellent agreement with the results reported in [13].

4.2 Density current

In this section we consider the classical density current benchmark proposed in [32]. The setup consists of a negative temperature perturbation in a motionless isentropic atmosphere with background potential temperature $\bar{\theta} = 300$ K and temperature

$$\bar{T} = \left(300 - zg\frac{\gamma R}{\gamma - 1}\right) \,\mathrm{K}$$



Figure 1: Inertia-gravity waves benchmark, potential temperature deviation a) t = 0 s, b) t = 3000 s. Contours are plotted from 0.001 K to 0.01 K with interval equal to 1×10^{-3} K for a) and from -0.0015 K to 0.003 K with interval equal to 5×10^{-4} K for b).

on the domain $\Omega = (-25.6, 25.6) \times (0, 6.4)$ km. More in detail, the temperature perturbation T' is defined as

$$T' = \begin{cases} 0 & \text{if } \tilde{r} > 1000\\ -15\frac{1+\cos(\pi\tilde{r})}{2} & \text{if } \tilde{r} \le 1000 \end{cases}$$
(20)

where $\tilde{r} = \sqrt{\left[\frac{(x-x_c)}{x_r}\right]^2 + \left[\frac{(x-y_c)}{y_r}\right]^2}$, $x_c = 0 \text{ m}$, $x_r = 4000 \text{ m}$, $y_c = 3000 \text{ m}$ and $x_r = 2000 \text{ m}$. Following [32], diffusive terms are included in order to stabilize the flow and, therefore, the classical Navier-Stokes equations in conservative form are considered. As explained in [23] and mentioned in Section 2, we resort to an operator splitting approach between the hyperbolic part and the diffusive terms, whose discretization is carried out by the implicit part of the IMEX scheme. This choice is commonly made in numerical methods for atmospherics dynamics, see e.g. [20, 31]. We consider the diffusion coefficient $\nu = 75 \text{ m}^2/\text{s}$



Figure 2: Inertia-gravity waves benchmark, potential temperature deviation at t = 3000 s along z = 5 km height.

and we set the thermal conductivity value κ such that the Prandtl number is taken to be Pr = 0.76. The boundary conditions are periodic on the left and right boundaries and wall boundary conditions on the top and bottom boundaries. The grid is composed by 1024×128 elements with r = 2 leading to a resolution equal to 25 m. The time step is taken equal to 0.1 s, yielding a maximum Courant number $C \approx 1.4$ and $C_u \approx 0.15$. Figure 3 shows the deviation of the potential temperature with respect to the background value at different times for the subdomain $(0, 19.2) \times (0, 4.8)$ km. In view of the negative buoyancy, the structure falls and reaches the bottom boundary. It then moves to the right, developing vortices. The front location is located at x = 15700 m, in agreement with the results obtained in [3, 21].

4.3 Cold bubble

In this Section, we consider a test case proposed in [26]. The computational domain is the rectangle $(0, 1000) \times (0, 2000)$ m and the initial condition is represented by a thermal anomaly introduced in an isentropic background atmosphere with constant potential temperature $\bar{\theta} = 303$ K. The perturbation of potential temperature θ' defines the initial datum and it is given by

$$\theta' = \begin{cases} A & \text{if } \tilde{r} \le r_0 \\ A \exp\left(-\frac{\left(\tilde{r} - r_0\right)^2}{\sigma^2}\right) & \text{if } \tilde{r} > r_0, \end{cases}$$
(21)

with $\tilde{r}^2 = (x - x_0)^2 + (z - z_0)^2$ and $x_0 = 500 \text{ m}$, $z_0 = 1250 \text{ m}$, $r_0 = 50 \text{ m}$, $\sigma = 100 \text{ m}$ and A = -15 K. The expression of the initial profile of the Exner pressure is given by

$$\overline{\pi} = 1 - \frac{g}{c_p \theta} z,$$

with $c_p = \frac{\gamma}{\gamma - 1}R = 1004.5 \,\mathrm{J\,kg^{-1}\,K^{-1}}$ denoting the specific heat at constant pressure. Notice that, unlike in [26], no artificial viscosity has been added to stabilize the computation. Wall boundary conditions are imposed at all the boundaries. The time step is taken to be $\Delta t = 0.08 \,\mathrm{s}$, corresponding to a maximum Courant number $C \approx 5.6$ and a maximum advective Courant number $C_u \approx 0.24$, whereas the final time is $T_f = 200 \,\mathrm{s}$. The computational grid



Figure 3: Current density benchmark, potential temperature deviation from background at: a) t = 0 s, b) t = 300 s, c) t = 600 s, d) t = 900 s. Contours are plotted from -16 K to -1 K with interval equal to 1 K.

is composed by 200×400 elements with r = 1 leading to a resolution equal to 5 m. Figure 4 shows the contours of the potential temperature deviation from the background at t = 0 s, t = 100 s and t = 200 s and the results are in reasonable agreement with those reported in [26].

We now demonstrate the h-adaptivity capabilities available in the proposed implementation. We use as refinement indicator the gradient of the potential temperature. More specifically, we set

$$\eta_K = \max_{i \in \mathcal{N}_K} |\nabla \theta|_i \tag{22}$$

where \mathcal{N}_K denotes the set of nodes over the element K. The initial computational grid is composed by 50×100 elements and only two local refinements are allowed, so as to control the advective Courant number and to match at the finest refinement level the resolution of the non adaptive mesh simulation. As one can easily notice from Figure 5, the refinement criterion is able to track the bubble. The contour plots in Figure 6 show a reasonable agreement for t = 50 s and t = 100 s, whereas for t = 200 s significant differences between the simulations with uniform and adaptive grid appear, which are due to the different development of the Kelvin-Helmholtz instability. The final mesh consists of 19334 elements instead of the 80000 elements of the full resolution mesh and a 25% reduction in computational time is achieved.

4.4 3D rising bubble

In this Section we consider the 3D rising bubble benchmark proposed in [21]. A neutrally stratified isentropic atmosphere is assumed, with $\bar{\theta} = 300$ K in the domain $\Omega = (-500, -500, -500) \times (500, 500, 1000)$ m. A spherical perturbation θ' located at $(x_0, y_0, z_0) = (0, 0, 350)$ m is added to the potential temperature

$$\theta' = \begin{cases} 0.25 \left[1 + \cos\left(\frac{\pi r}{r_0}\right) \right] & \text{if } \tilde{r} \le r_0 \\ 0 & \text{if } \tilde{r} > r_0, \end{cases}$$
(23)





Figure 4: Cold bubble benchmark, potential temperature deviation from background at: a) t = 0 s, b) t = 100 s, c) t = 200 s. Contours are plotted from -15 K to -1 K with interval equal to 1 K for a) and from -11.5 K to 8 K with interval equal to 1.625 K for b) and c).

with $\tilde{r} = \sqrt{(x - x_0)^2 + (y - y_0)^2 + (z - z_0)^2}$ and $r_0 = 250$ m. Wall boundary conditions are imposed for all the six boundaries and we take r = 2. In order to enhance the computational efficiency, we use the *h*-adaptivity capabilities, with the same refinement indicator (22) introduced in the previous Section. The initial grid is composed by $24 \times 24 \times 36$ elements and we allowed up to two local refinements which would correspond to a uniform mesh with $96 \times 96 \times 144$ elements and to a resolution around 5 m. The time step is taken to be equal to $\Delta t = 0.4$ s, leading to a maximum acoustic Courant number $C \approx 27$ and advective Courant number $C_u \approx 0.22$. Figure 7 shows snapshots of the bubble at t = 200 s and t = 400 s. At the later time, a Kelvin-Helmholtz instability starts to develop, which is however still insufficiently well resolved by the present mesh. Further refinement levels or higher polynomial degrees will have to be employed in future simulations to achieve better accuracy at the later stage. Similar issue for an analogous test case are reported in [6].



Figure 5: Cold bubble benchmark, adaptive mesh at: a) t = 100 s, b) t = 200 s.

The final grid is composed by 62792 elements.

The size of this benchmark makes it a good candidate for a parallel scaling test. An initial mesh composed by $48 \times 48 \times 72$ elements corresponding to 13436928 dofs for the velocity and 4478976 dofs for the remaining scalar variables is considered. Two configurations are employed: in the first case we keep it fixed, whereas in the second one we apply h-adaptivity with two local refinements, roughly doubling the number of degrees of freedom. A strong scaling analysis is performed executing the simulation up to time t = 8 s and we use from 32 up to 1024 2xCPU x86 Intel Xeon Platinum 8276-8276L @ 2.4Ghz cores of the HPC infrastructure GALILEO100 at the Italian supercomputing center CINECA.

The results, reported in Figure 8, are quite similar for the two configurations. A good linear scaling is obtained up to 128 cores, even with superlinear behaviour for the fixed mesh framework due to cache effects. Starting from 256 cores, the performance of the fixed mesh configuration exhibits a small degradation and, for a higher number of cores, the speed-up is less optimal for both configurations, due to overwhelming communication costs. The apparent better behaviour of the h-adaptive version is due to the fact that more degrees of freedoms are involved and, therefore, the role of communication costs is less evident. The result also highlights that the local refining procedure has no significant impact on the parallel performances and that both efficiency and scalability can be achieved in this framework. These results are also similar to those obtained in [24] for a fixed grid with an implicit incompressible solver built in the same framework and for which the employed method coincides with the implicit part of the IMEX scheme presented in Section 3.



Figure 6: Cold bubble benchmark, potential temperature deviation from background at: a) t = 50 s, b) t = 100 s, c) t = 200 s. Black line denotes the results with uniform grid, whereas red line represents the results with *h*-adaptivity. Contours are plotted from -11.5 to 8 with interval equal to 1.625 K.

5 Numerical results with orography

We now consider a number of tests concerning idealized flows over orography, that since the seminal papers [17, 18] have become a standard benchmark for numerical models of atmospheric flows, see e.g. the results and discussions in [4, 21, 25, 34].

5.1 Hydrostatic flow over a hill

In a first test, the bottom boundary is described by the function

$$h(x) = \frac{h_m}{1 + \left(\frac{x - x_c}{a_c}\right)^2},\tag{24}$$



Figure 7: 3D rising bubble benchmark, results for y = 0 m; on the left the adaptive meshes at t = 200 s (a) and t = 400 s (c) are reported, whereas on the right potential temperature deviations from the background at t = 200 s (b) and t = 400 s (d) are reported.

the so-called versiera di Agnesi, where h_m is the height of the hill and a_c is the halfwidth. The classical Gal-Chen height-based terrain-following coordinate [10] is used to build the mapping between the reference element and the physical one and to obtain a terrain following mesh in Cartesian coordinates. We first consider the linear hydrostatic configuration presented e.g. in [13]. The computational domain is $\Omega = (0, 240) \times (0, 30)$ km with $h_m = 1 \text{ m}, x_c = 120 \text{ km}$ and $a_c = 10 \text{ km}$. The final time is $T_f = 45000 \text{ s}$. The initial state of the atmosphere consists of a constant mean flow with $\overline{u} = 20 \text{ m s}^{-1}$ and of an isothermal background profile with temperature $\overline{T} = 250 \text{ K}$. The initial profile of the



Figure 8: 3D rising bubble benchmark, strong scaling analysis. The speedup is computed with respect to the time required with 32 cores.

Exner pressure is given by

$$\overline{\pi} = \left(\frac{p_0}{\overline{p}}\right)^{\frac{\gamma-1}{\gamma}} = \exp\left(-\frac{g}{c_p\overline{T}}z\right).$$

We recall that $c_p = \frac{\gamma}{\gamma - 1}R$ denotes the specific heat at constant pressure and that here $p_0 = 10^5$ Pa; moreover, since in an isothermal configuration the Brunt-Väisälä frequency is $N = \frac{g}{\sqrt{c_p T}}$, it can be easily checked that $\frac{Na_c}{\overline{u}} >> 1$, so that this configuration corresponds to a hydrostatic regime according to the classification in [25]. For what concerns the boundary conditions, wall boundary conditions are used for the bottom boundary and non-reflecting boundary conditions are required by the top boundary and the lateral boundaries. For this purpose, we introduce a Rayleigh damping profile following [21]

$$\lambda = \begin{cases} 0, & \text{if } z < z_B \\ \overline{\lambda} \sin^2 \left[\frac{\pi}{2} \left(\frac{z - z_B}{z - z_T} \right) \right] & \text{if } z \ge z_B, \end{cases}$$
(25)

where z_B denotes the height at which the damping starts and z_T is the top height of the considered domain. Analogous definitions apply for the two lateral boundaries. In this case, we consider $\overline{\lambda}\Delta t = 0.3$ and we apply the damping layer in the topmost 15 km of the domain and in the first and last 80 km along the horizontal direction. The grid is composed by 100×75 elements with r = 4, yielding a resolution of 600 m along x and 100 m along z, whereas the time-step is equal to 2.5 s, leading to $C \approx 1.84$ and $C_u \approx 0.12$. Following [29], we also define the vertical momentum flux as

$$m(z) = \int_{-\infty}^{\infty} \overline{\rho}(z) u'(x, z) w'(x, z) dx, \qquad (26)$$

where u' and w' represent the deviation from the background state of the horizontal and vertical velocity, respectively. This is a very important diagnostic quantity in atmospheric modelling, used to check that the numerical model is correctly reproducing the orographic forcing on the main flow. From the linear theory, the analytical momentum flux is given by

$$m^{H} = -\frac{\pi}{4} \overline{\rho}_{s} \overline{u}_{s} N h_{m}^{2}, \qquad (27)$$

where $\overline{\rho}_s$ and \overline{u}_s denote the surface background density and velocity, respectively. Figure 9 shows the behaviour over time of the momentum flux normalized by its analytical value. It can be noticed that the analytical value is approached as the simulation reaches the steady state.



Figure 9: Linear hydrostatic flow over a hill, evolution of normalized momentum flux.

We consider now the more challenging, nonlinear hydrostatic case considered in [4, 25]. The computational domain is $\Omega = (0, 512) \times (0, 28)$ km with $h_m = 800 \text{ m}, x_c = 256 \text{ km}$ and $a_c = 16 \text{ km}$. The final time is $T_f = 60\,000 \text{ s}$. The damping layer is applied starting from z = 11.5 km and in the first and last 172 km along the horizontal direction. The background velocity is $\overline{u} = 32 \text{ m s}^{-1}$ and the Brunt-Väisälä frequency N is equal to 0.02 s^{-1} . The mesh is composed by 160×112 elements with r = 2, yielding a resolution of 1600 m along x and 125 m along z, whereas the time step is equal to 10 s, yielding a maximum Courant number $C \approx 1.41$ and $C_u \approx 0.25$. Figure 11 shows the contour plots of both the horizontal velocity perturbation and vertical velocity, which compare well with those presented e.g. in [25]. The behaviour over time of the normalized momentum flux is reported in Figure 12 and its value at the surface at $t = T_f$ is approximately equal to 1.22, which is comparable to the one obtained in [25]. The momentum flux differs from the analytical one because we are no more in a linear regime. Moreover, as explained in [9], it is strongly dependent on the position of the absorbing layer. These results confirm the stability and the accuracy of the proposed numerical scheme also in presence of orography.

5.2 Nonhydrostatic flow over a hill

In this section we consider the nonhydrostatic regime, characterized by $\frac{Na_c}{\overline{u}} \approx 1$. The bottom boundary is again described by the function (24). We first adopt the linear nonhydrostatic configuration described e.g. in [13]. The computational domain is $\Omega = (0, 144) \times (0, 30)$ km with $h_m = 1 \text{ m}, x_c = 72 \text{ km}$ and $a_c = 1 \text{ km}$. The final time is $T_f = 28800 \text{ s}$. The initial state of the atmosphere is described by the following potential temperature and Exner pressure, respectively:

$$\overline{\theta} = \theta_{ref} \exp\left(\frac{N^2}{g}z\right) \tag{28}$$

$$\overline{\pi} = 1 + \frac{g^2}{c_p \theta_{ref} N^2} \left[\exp\left(-\frac{N^2}{g}z\right) - 1 \right],\tag{29}$$



Figure 10: Linear hydrostatic flow over a hill at $t = T_f$, a) horizontal velocity deviation. Contour values are between $-2.5 \cdot 10^{-2} \,\mathrm{m \, s^{-1}}$ and $2.5 \cdot 10^{-2} \,\mathrm{m \, s^{-1}}$ with an interval equal to $5 \cdot 10^{-3} \,\mathrm{m \, s^{-1}}$, b) vertical velocity. Contour values are between $-4 \cdot 10^{-3} \,\mathrm{m \, s^{-1}}$ and $4 \cdot 10^{-3} \,\mathrm{m \, s^{-1}}$ with an interval equal to $5 \cdot 10^{-4} \,\mathrm{m \, s^{-1}}$.

with $\theta_{ref} = 280 \,\mathrm{K}$ and $N = 0.01 \,\mathrm{s}^{-1}$. The background velocity \overline{u} is equal to $10 \,\mathrm{m} \,\mathrm{s}^{-1}$. Following [17], the analytical momentum flux is given by

$$m^{NH} = 0.457m^H \tag{30}$$

and this value will be used to compute the normalized momentum flux for the present case. Wall boundary conditions are applied on the bottom boundary and non-reflecting boundary conditions are employed on the top and lateral boundaries with $\overline{\lambda}$ such that $\overline{\lambda}\Delta t = 0.15$. The damping layer is applied in the topmost 14 km of the domain and in the first and last 40 km along the horizontal direction. The mesh is composed by 200 × 50 elements with r = 4, yielding a resolution of 180 m along x and 150 m along z, whereas the time step is equal to 1 s, leading to $C \approx 2.02$ and $C_u \approx 0.06$. Figure 13 reports the



Figure 11: Nonlinear hydrostatic flow over a hill at $t = T_f$, a) horizontal velocity deviation, values between $-23 \,\mathrm{m \, s^{-1}}$ and $28 \,\mathrm{m \, s^{-1}}$ with contour interval of $2 \,\mathrm{m \, s^{-1}}$, b) vertical velocity, values between $-3.9 \,\mathrm{m \, s^{-1}}$ and $3.5 \,\mathrm{m \, s^{-1}}$ with contour interval of $0.5 \,\mathrm{m \, s^{-1}}$.

time evolution of the normalized momentum flux and, as for the linear hydrostatic case in Section 5.1, the analytical value is approached as the simulation reaches the steady state.

Finally, we consider a nonlinear nonhydrostatic case, see e.g. [34]. The computational domain is $\Omega = (0, 40) \times (0, 20)$ km with $h_m = 450 \text{ m}, x_c = 20 \text{ km}, a_c = 1 \text{ km}, T_f = 36\,000 \text{ s}, N = 0.02 \text{ s}^{-1}, \theta_{ref} = 273 \text{ K}$ and $\overline{u} = 13.28 \text{ m s}^{-1}$. The damping layer is applied in the topmost 11 km of the domain and in the first and last 10 km along the horizontal direction. The mesh is composed by 50×50 elements with r = 4, yielding a resolution of 200 m along x and 100 m along z. The time step is equal to 0.5 s, leading to a maximum Courant number $C \approx 1.13$ and $C_u \approx 0.08$. Figure 14 shows the contour plots of both horizontal velocity perturbation and vertical velocity, which are analogous to those reported in [34], as well as the time evolution of the normalized vertical momentum flux. Notice that the momentum flux is normalized by the analytical value (27).



Figure 12: Nonlinear hydrostatic flow over a hill, normalized momentum flux evolution



Figure 13: Linear nonhydrostatic flow over a hill, evolution of normalized momentum flux.

5.3 Schär hill

In this Section, we consider the well known Schär mountain test, which consists of a steadystate flow over a five-peak mountain chain [21, 27]. The domain is $\Omega = (-50, 50) \times (0, 30)$ km with surface temperature $T_{ref} = 288$ K, constant buoyancy frequency $N = 0.01 \text{ s}^{-1}$ and a background wind $\bar{u} = 10 \text{ m s}^{-1}$. The mountain profile is defined as

$$h(x) = h_m \exp\left[\left(-\frac{x}{a_c}\right)^2\right] \cos^2\left(\frac{\pi x}{\lambda_c}\right),\tag{31}$$

with $h_m = 250 \text{ m}, a_c = 5 \text{ km}$ and $\lambda_c = 4 \text{ km}$. The background density and pressure have the same expression as in (17)-(18) with $\theta_{ref} = 288 \text{ K}$ and the final time is $T_f = 18000 \text{ s}$. The damping layer is applied in the topmost 10 km of the domain and in the first and last 10 km along the horizontal direction with $\overline{\lambda}\Delta t = 1.2$. The mesh is composed by 100×50



Figure 14: Nonlinear nonhydrostatic flow over a hill, a) horizontal velocity deviation, values between $-7.2 \,\mathrm{m \, s^{-1}}$ and $9.0 \,\mathrm{m \, s^{-1}}$ with contour interval of $1.16 \,\mathrm{m \, s^{-1}}$ at $t = T_f$, b) vertical velocity, values between $-4.2 \,\mathrm{m \, s^{-1}}$ and $4.0 \,\mathrm{m \, s^{-1}}$ with contour interval of $0.586 \,\mathrm{m \, s^{-1}}$ at $t = 18\,000 \,\mathrm{s}$, c) normalized momentum flux evolution.

elements with r = 4, leading to a resolution of 250 m along x and of 150 m along z, whereas the time step is equal to 2.5 s, yielding a maximum acoustic Courant number $C \approx 2.02$ and a maximum advective Courant number $C_u \approx 0.09$. Figure 15 shows the contour plots of both horizontal velocity perturbation and vertical velocity, which are analogous to those reported in [13, 21].

5.4 3D medium-steep bell-shaped hill

In this Section, we consider the three-dimensional flow over a bell-shaped hill, see e.g. [21]. The computation domain is $\Omega = (0, 60) \times (0, 40) \times (0, 20)$ km. The mountain profile is defined as

$$h(x,y) = \frac{h_m}{\left[1 + \left(\frac{x - x_c}{a_c}\right)^2 + \left(\frac{y - y_c}{a_c}\right)^2\right]^{3/2}},$$
(32)

with $h_m = 400 \text{ m}, a_c = 1 \text{ km}, x_c = 30 \text{ km}$ and $y_c = 20 \text{ km}$. We consider as buoyancy frequency $N = 0.01 \text{ s}^{-1}$ and a background velocity $\overline{u} = 10 \text{ m s}^{-1}$. We are therefore in a nonhydrostatic regime since $\frac{Na_c}{\overline{u}} = 1$. The background density and pressure have the same expression of (17) and (18) with $\theta_{ref} = 293.15 \text{ K}$ and the final time is $T_f = 3600 \text{ s}$. The damping layer is applied in the topmost 6 km of the domain and in the first and last 20 km along the lateral boundaries with $\overline{\lambda}\Delta t = 1.2$. In order to increase the resolution around



Figure 15: Schär mountain test case at $t = T_f$, a) horizontal velocity deviation. Contour values are between $-2 \,\mathrm{m \, s^{-1}}$ and $2 \,\mathrm{m \, s^{-1}}$ with an interval equal to $0.2 \,\mathrm{m \, s^{-1}}$, b) vertical velocity. Contour values are between $-0.5 \,\mathrm{m \, s^{-1}}$ and $0.5 \,\mathrm{m \, s^{-1}}$ with an interval equal to $5 \cdot 10^{-2} \,\mathrm{m \, s^{-1}}$.

the hill, we consider a non-uniform grid by taking a resolution of 250 m between x = 25 km and x = 40 km and a resolution of 250 m between y = 12 km and y = 28 km. A uniform resolution of 500 m is considered along the vertical direction z, as well as for the remaining part of the lateral boundaries. A snapshot of the grid for the x - y sloice at z = 800 m is reported in Figure 16. The mesh is composed by 8288 elements with polynomial degree r = 4, whereas the time step is equal to 2 s, yielding a maximum acoustic Courant number $C \approx 1.95$ and a maximum advective Courant number $C_u \approx 0.1$. The results in Figure 17 are in agreement with those reported in [21].



Figure 16: 3D mountain benchmark, x-y slice at z = 800 m of the computational grid.

6 Conclusions

We have presented an application of semi-implicit discretization of the compressible Euler equations proposed in [23] to atmospheric dynamics. The method includes non-conforming h— refinement, as implemented in the framework provided by the numerical library *deal.II* [1, 2]. Being based on an open source library enhances code portability and embeds the development of codes specific for numerical weather prediction applications in the framework of a development project that is being continuously adapted to novel architectures and extended to include novel versions of the DG method. Simulation of classical benchmarks demonstrates that the method can simulate accurately small scale flows in presence of gravity and idealized flows over orography. In future developments, we aim to show the potential of non-conforming mesh refinement to increase locally the resolution over complex orography. Moreover, we will consider the inclusion of more complex physical phenomena, such as for example turbulence, water vapour transport and adiabatic heating, in order to demonstrate that all the typical features of a high resolution numerical weather prediction model can be included in the proposed adaptive framework without significant loss of accuracy or efficiency.

Acknowledgements

The simulations have been partly run at CINECA thanks to the computational resources made available through the ISCRA-C projects SIDICoNS - HP10CLPLXI and NUMNETF - HP10C06Y02. This work has been partly supported by the ESCAPE-2 project, European Union's Horizon 2020 Research and Innovation Programme (Grant Agreement No. 800897).



Figure 17: 3D mountain benchmark, x-y slice at z = 800 m of the vertical velocity. The values are between -1.5 m s^{-1} and 1.3 m s^{-1} with contour interval of 0.1 m s^{-1}

References

- D. Arndt, W. Bangerth, M. Feder, M. Fehling, R. Gassmöller, T. Heister, L. Heltai, M. Kronbichler, M. Maier, P. Munch, J.-P. Pelteret, S. Sticko, B. Turcksin, and D. Wells. The deal II library, version 9.4. *Journal of Numerical Mathematics*, 30(3):231–246, 2022.
- [2] W. Bangerth, R. Hartmann, and G. Kanschat. deal II: a general-purpose objectoriented finite element library. ACM Transactions on Mathematical Software, 33:24– 51, 2007.
- [3] T. Benacchio, W.P. O'Neill, and R. Klein. A blended soundproof-to-compressible numerical model for small- to mesoscale atmospheric dynamics. *Monthly Weather Review*, 142:4416–4438, 2014.
- [4] L. Bonaventura. A Semi-Implicit, Semi-Lagrangian Scheme Using the Height Coordinate for a Nonhydrostatic and Fully Elastic Model of Atmospheric Flows. *Journal of Computational Physics*, 158:186–213, 2000.
- [5] L. Bonaventura, R. Redler, and R. Budich. Earth System Modelling 2: Algorithms, Code Infrastructure and Optimisation. Springer-Verlag, 2012.
- [6] S. Busto, M. Tavelli, W. Boscheri, and M. Dumbser. Efficient high order accurate staggered semi-implicit discontinuous Galerkin methods for natural convection problems. *Computers & Fluids*, 198:104399, 2020.
- [7] V. Casulli and D. Greenspan. Pressure method for the numerical solution of transient, compressible fluid flows. *International Journal for Numerical Methods in Fluids*, 4:1001–1012, 1984.

- [8] M. Dumbser and V. Casulli. A conservative, weakly nonlinear semi-implicit finite volume scheme for the compressible Navier-Stokes equations with general equation of state. Applied Mathematics and Computation, 272:479–497, 2016.
- [9] D. R. Durran and J. B. Klemp. A compressible model for the simulation of moist mountain waves. *Monthly Weather Review*, 111:2341–2361, 1983.
- [10] T. Gal-Chen and R.C.J. Somerville. On the use of a coordinate transformation for the solution of the Navier-Stokes equations. *Journal of Computational Physics*, 17:209– 228, 1975.
- [11] F.X. Giraldo. Semi-implicit time-integrators for a scalable spectral element atmospheric model. Quarterly Journal of the Royal Meteorological Society, 131:2431–2454, 2005.
- [12] F.X. Giraldo, J.F. Kelly, and E.M. Constantinescu. Implicit-Explicit formulations of a three-dimensional nonhydrostatic unified model of the atmosphere (NUMA). SIAM Journal of Scientific Computing, 35:1162 – 1194, 2013.
- [13] F.X. Giraldo and M. Restelli. A study of spectral element and discontinuous Galerkin methods for the Navier-Stokes equations in nonhydrostatic mesoscale atmospheric modeling: Equation sets and test cases. *Journal of Computational Physics*, 227:3849– 3877, 2008.
- [14] F.X. Giraldo, M. Restelli, and M. Läuter. Semi-implicit formulations of the Navier-Stokes equations: Application to nonhydrostatic atmospheric modeling. *SIAM Journal* on Scientific Computing, 32:3394–3425, 2010.
- [15] J. F. Kelly and F. X. Giraldo. Continuous and discontinuous Galerkin methods for a scalable three-dimensional nonhydrostatic atmospheric model: Limited-area mode. *Journal of Computational Physics*, 231(24):7988–8008, 2012.
- [16] C.A. Kennedy and M.H. Carpenter. Additive Runge-Kutta schemes for convectiondiffusion-reaction equations. *Applied Numerical Mathematics*, 44:139 – 181, 2003.
- [17] J.B. Klemp and D.R. Durran. An upper boundary condition permitting internal gravity wave radiation in numerical mesoscale models. *Monthly Weather Review*, 111:430–444, 1983.
- [18] J.B. Klemp and D.K. Lilly. Numerical simulation of hydrostatic mountain waves. Journal of the Atmospheric Sciences, 35:78–107, 1978.
- [19] M.A. Kopera and F.X. Giraldo. Analysis of adaptive mesh refinement for IMEX Discontinuous Galerkin solutions of the compressible Euler equations with application to atmospheric simulations. *Journal of Computational Physics*, 275:92–117, 2014.
- [20] C. Kühnlein, W. Deconinck, R. Klein, S. Malardel, Z.P. Piotrowski, P.K. Smolarkiewicz, J. Szmelter, and N.P. Wedi. FVM 1.0: A nonhydrostatic finite-volume dynamical core formulation for IFS. *Geoscientific Model Development*, 12:651–676, 2019.
- [21] T. Melvin, T. Benacchio, B. Shipway, N. Wood, J. Thuburn, and C.J. Cotter. A mixed finite-element, finite-volume, semi-implicit discretization for atmospheric dynamics: Cartesian geometry. *Quarterly Journal of the Royal Meteorological Society*, 145:2835– 2853, 2019.
- [22] A. Müller, J. Behrens, F.X. Giraldo, and V. Wirth. Comparison between adaptive and uniform Discontinuous Galerkin simulations in dry 2D bubble experiments. *Journal* of Computational Physics, 235:371–393, 2013.
- [23] G. Orlando, P. F. Barbante, and L. Bonaventura. An efficient IMEX-DG solver for the compressible Navier-Stokes equations for non-ideal gases. *Journal of Computational Physics*, 471:111653, 2022.

- [24] G. Orlando, A. Della Rocca, P. F. Barbante, L. Bonaventura, and N. Parolini. An efficient and accurate implicit DG solver for the incompressible Navier-Stokes equations. *International Journal for Numerical Methods in Fluids*, 94:1484–1516, 2022.
- [25] J.P. Pinty, R. Benoit, E. Richard, and R. Laprise. Simple tests of a semi-implicit semi-Lagrangian model on 2D mountain wave problems. *Monthly Weather Review*, 123:3042–3058, 1995.
- [26] M. Restelli and F.X. Giraldo. A conservative Discontinuous Galerkin semi-implicit formulation for the Navier-Stokes equations in nonhydrostatic mesoscale modeling. *SIAM Journal of Scientific Computing*, 31:2231–2257, 2009.
- [27] C. M. Schär, D. Leuenberger, O. Fuhrer, D. Lüthi, and C. Girard. A new terrainfollowing vertical coordinate formulation for atmospheric prediction models. *Monthly Weather Review*, 130:2459–2480, 2002.
- [28] W.C. Skamarock and J. B. Klemp. Efficiency and accuracy of the Klemp-Wilhelmson time-splitting technique. *Monthly Weather Review*, 122:2623–2630, 1994.
- [29] R.B. Smith. The influence of mountains on the atmosphere. Advances in Geophysics, 21:87–230, 1979.
- [30] P.K. Smolarkiewicz, C. Kühnlein, and N.P. Wedi. Semi-implicit integrations of perturbation equations for all-scale atmospheric dynamics. *Journal of Computational Physics*, 376:145–159, 2019.
- [31] J. Steppeler, R. Hess, G. Doms, U. Schättler, and L. Bonaventura. Review of numerical methods for nonhydrostatic weather prediction models. *Meteorology and Atmospheric Physics*, 82:287–301, 2003.
- [32] J. Straka, R. Wilhelmson, L.J. Wicker, J. Anderson, and K. Droegemeier. Numerical solutions of a non-linear density current: A benchmark solution and comparisons. *International Journal for Numerical Methods in Fluids*, 17:1–22, 1993.
- [33] G. Strang. On the construction and comparison of difference schemes. SIAM Journal on Numerical Analysis, 5:506–517, 1968.
- [34] G. Tumolo and L. Bonaventura. A semi-implicit, semi-Lagrangian discontinuous Galerkin framework for adaptive numerical weather prediction. *Quarterly Journal* of the Royal Meteorological Society, 141:2582–2601, 2015.
- [35] L. Yelash, A. Müller, M. Lukáčová-Medvid'ová, F.X. Giraldo, and V. Wirth. Adaptive discontinuous evolution Galerkin method for dry atmospheric flow. *Journal of Computational Physics*, 268:106–133, 2014.

MOX Technical Reports, last issues

Dipartimento di Matematica Politecnico di Milano, Via Bonardi 9 - 20133 Milano (Italy)

- **67/2022** Alghamdi, M. M.; Boffi, D.; Bonizzoni, F. *A greedy MOR method for the tracking of eigensolutions to parametric elliptic PDEs*
- 64/2022 Massi, M.C., Dominoni, L., Ieva, F., Fiorito, G. *A Deep Survival EWAS approach estimating risk profile based on pre-diagnostic DNA methylation: an application to Breast Cancer time to diagnosis*
- **63/2022** Corti, M.; Antonietti, P.F.; Dede', L.; Quarteroni, A. Numerical Modelling of the Brain Poromechanics by High-Order Discontinuous Galerkin Methods
- **65/2022** Dassi, F.; Fumagalli, A.; Mazzieri, I.; Vacca, G. *Mixed Virtual Element approximation of linear acoustic wave equation*
- **66/2022** Antonietti, P.F.; Liverani, L.; Pata, V. Lack of superstable trajectories in linear viscoelasticity: A numerical approach
- 62/2022 Ciaramella, G.; Halpern, L.; Mechelli, L. Convergence analysis and optimization of a Robin Schwarz waveform relaxation method for periodic parabolic optimal control problems
- **61/2022** Gregorio, C.; Cappelletto, C.; Romani, S.; Stolfo, D.; Merlo, M.; Barbati, G. Using marginal structural joint models to estimate the effect of a time-varying treatment on recurrent events and survival: An application on arrhythmogenic cardiomyopathy
- **59/2022** Boon, W. M.; Fumagalli, A. A multipoint vorticity mixed finite element method for incompressible Stokes flow
- **60/2022** Cortellessa, D.; Ferro, N.; Perotto, S.; Micheletti, S. Enhancing level set-based topology optimization with anisotropic graded meshes
- 58/2022 Zingaro, A.; Bucelli, M.; Fumagalli, I.; Dede', L; Quarteroni, A.Modeling isovolumetric phases in cardiac flows by an Augmented Resistive Immersed Implicit Surface Method